



Deep Percolation Characteristics via Field Moisture Sensor Measurements in Rice Experimental Field, Phitsanulok, Thailand

**Nittaya Kangboonma
Chulalongkorn University**

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Straightforward Overview

- I. Introduction
- II. Objectives and study area conditions
- III. Study Procedures and Theories used
- IV. Field investigations
- V. Results
- VI. Conclusions

In recent several years, surface water trends to be deficit in some areas. The insufficient surface water resources lead to increase groundwater abstraction especially in dry season. In this last ten years, the tendency of groundwater level declined in the Lower Yom and Nan River due to groundwater use mainly for irrigation use. Groundwater recharge mechanism knowledge is essential to understand groundwater movement especially in the area where groundwater use is excessive.

Deep percolation from irrigation water plays a key role in irrigation demand and groundwater supply by replenishing shallow aquifers at the local and regional scales. Percolation is accounted by vertical flow during the experiment

In the study, the soil moisture sensor module is used to detect the moisture of the soil or judge if there is water around the sensor. If more water is presented in the soil, the sensor would output resistance of soil which could covert to soil moisture. The motivation of soil moisture sensor is adapted from low-cost soil moisture profile probe [6]. Percolation identified water content in the soil flows below root zone efficiently [7]. The advantage of percolation is to induce recharge to groundwater and dilute chemicals in the soil. Hillel [8] concluded that percolation rate relied on both soil property and water content. Soil properties, affected to percolation, are porosity, void distribution, and void shape while flow properties, affected to percolation, are density and viscosity.

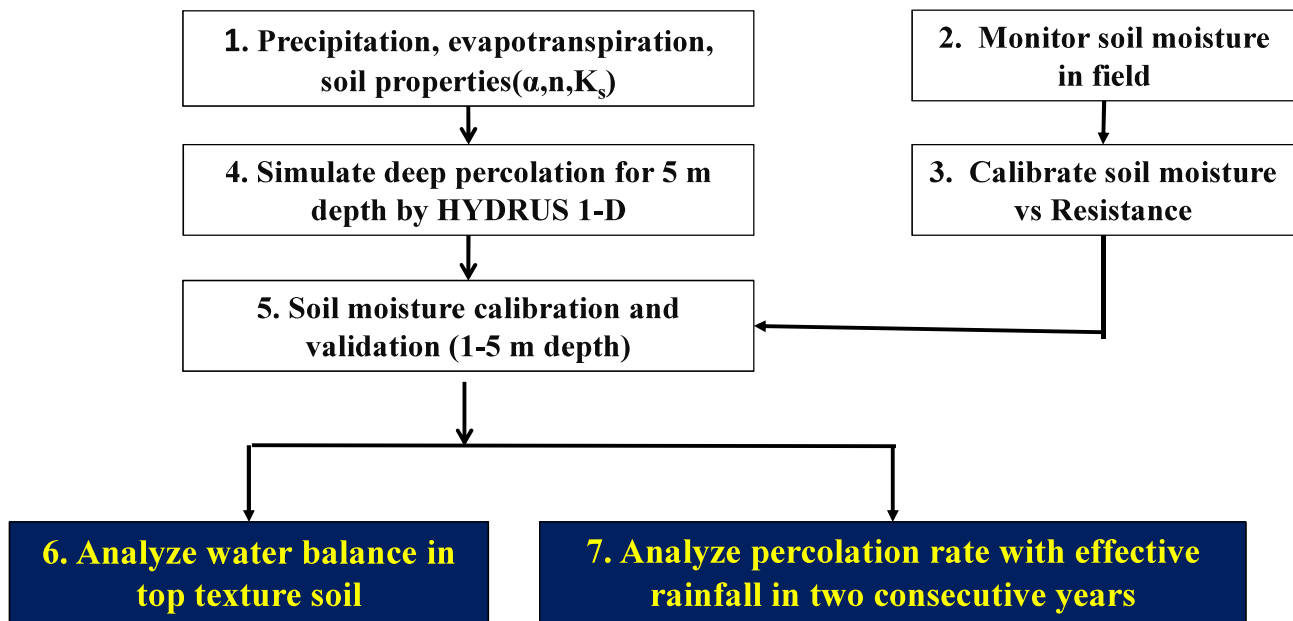
II Objectives

- This investigation aims to understand deep percolation mechanism (soil moisture change, water balance, percolation rate) from rainfall utilizing **field soil moisture monitoring approach**. Henceforth, the deeper percolation procedure and results will be useful for further determining groundwater yields and groundwater recharge in the consecutive drought years.

The study area

- Phitsnulok Province (north of center plain),
- A main irrigation area,
- Does not enough surface water in dry season.
- Conjunctive groundwater use.



Analysis groundwater and surface water interaction process

The governing flow equation for the uniform Darcian flow of water in a porous medium is adopted by the following modified form of the Richards' equation: (Simunek, Van Genuchten, and Sejna 2005)

$$\frac{\partial \theta}{\partial t} = \frac{\partial}{\partial x_i} \left[K \left(K_{ij}^A \frac{\partial h}{\partial x_j} + K_{iz}^A \right) \right] - S \quad (1)$$

θ is the volumetric water content, (L^3L^{-3})

K is the hydraulic conductivity (LT^{-1}),

h is the pressures head (L)

S is a sink term [T^{-1}]

x_i ($i=1,2$) are the spatial coordinates [L],

t is the time (T) and

z is the vertical ordinate (L)

K_{ij}^A are components of a dimensionless anisotropy tensor \mathbf{K}^A

K is the unsaturated hydraulic conductivity function [LT^{-1}] given by

$$K(h) = K_s S_e^{1/2} \left[1 - (1 - S_e^{1/m})^m \right]^2 \quad (2)$$

$$\theta(h) = \begin{cases} \theta_r + \frac{\theta_s - \theta_r}{[1 - |\alpha h|^n]^m} & h < 0 \\ \theta_s & h \geq 0 \end{cases} \quad (3)$$

$$S_e = \frac{\theta - \theta_r}{\theta_s - \theta_r} \quad (4)$$

$$m = 1 - 1/n, \quad n > 1 \quad (5)$$

S_e is the effective water content

θ_r denote the residual water content

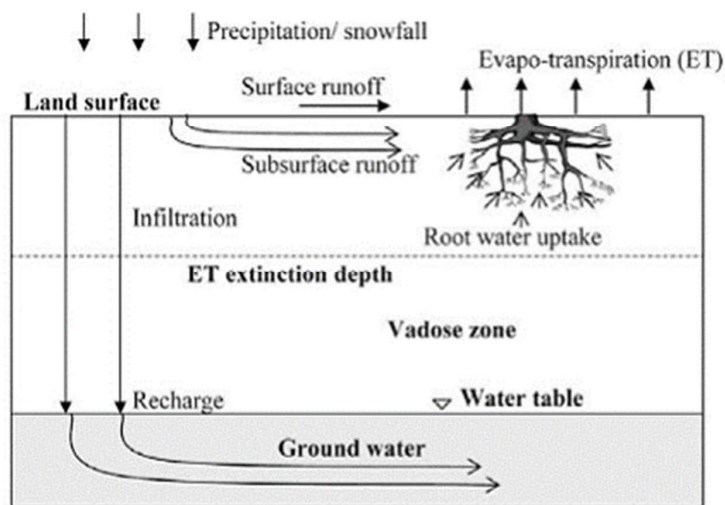
θ_s denote the saturated water content

K_s is the saturated hydraulic conductivity

α is the inverse of the air-entry value (or bubbling pressure)

n is a pore-size distribution index

III. STUDY Procedures and Theories used



A schematic showing the processes affecting subsurface hydrology (Source : Twarakaviet al., 2008)

Percolation

❑ Purpose

to determine land recharge rate and land recharge coefficient via soil moisture sensor

❑ Location

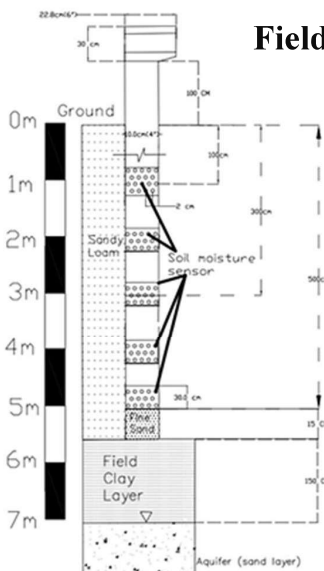
The experiment site was located in the Rice Water Use Experimental Station 2 of Royal Irrigation Department in Phitsanulok Province

❑ Study period

Daily (**350** days, July 2017-June 2018) measured soil moisture data in the field at 1-5 m depths

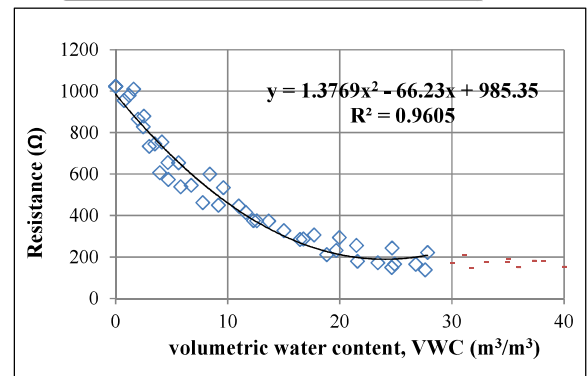
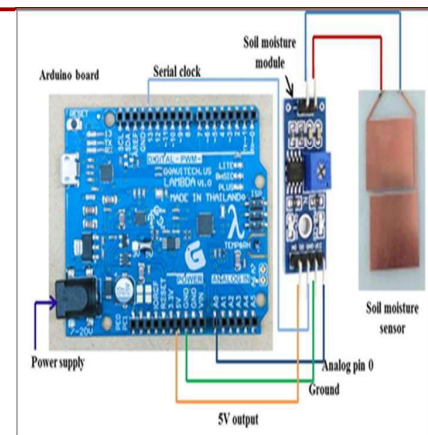
Field equipment installation

Field Data analysis



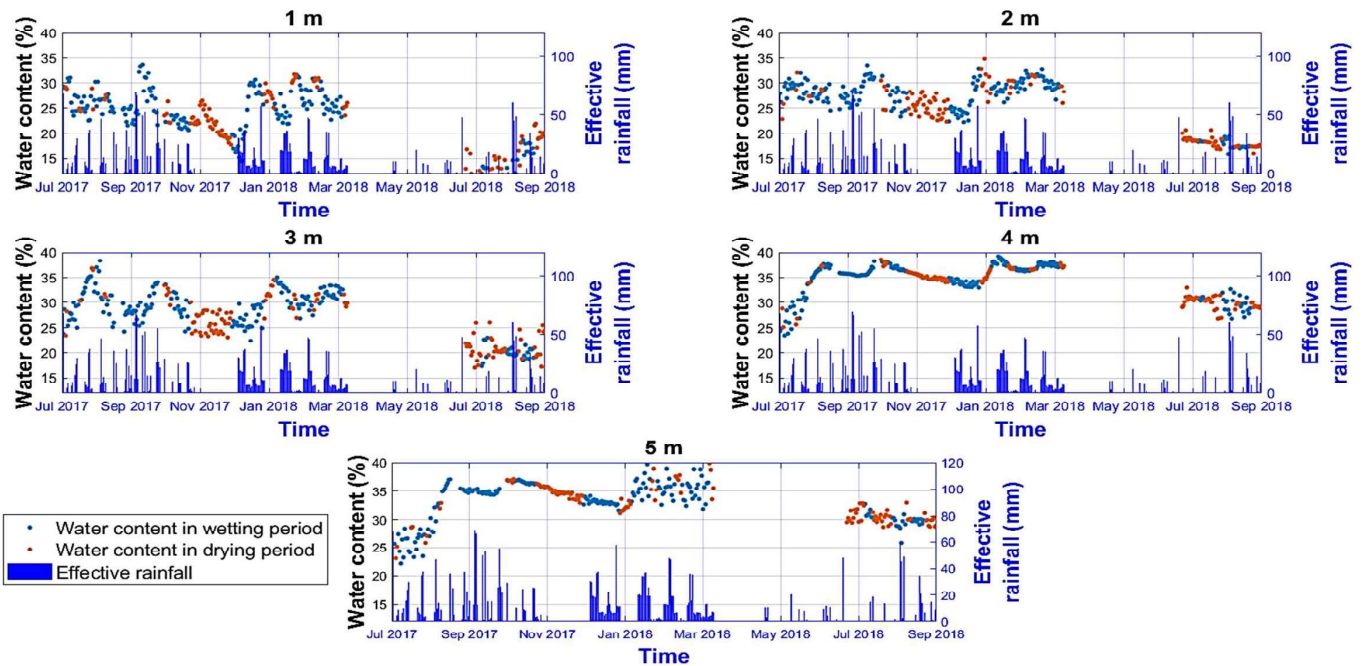
- Sandy Loam (0-4 m)
- Sandy Clay(5 m)

- Daily measurement (**350** days) soil moisture 1-5 m. depth
- For HYDRUS-1D simulation
- Inputs from field data
 - Rainfall
 - Evaporation
 - Transpiration
 - Groundwater level

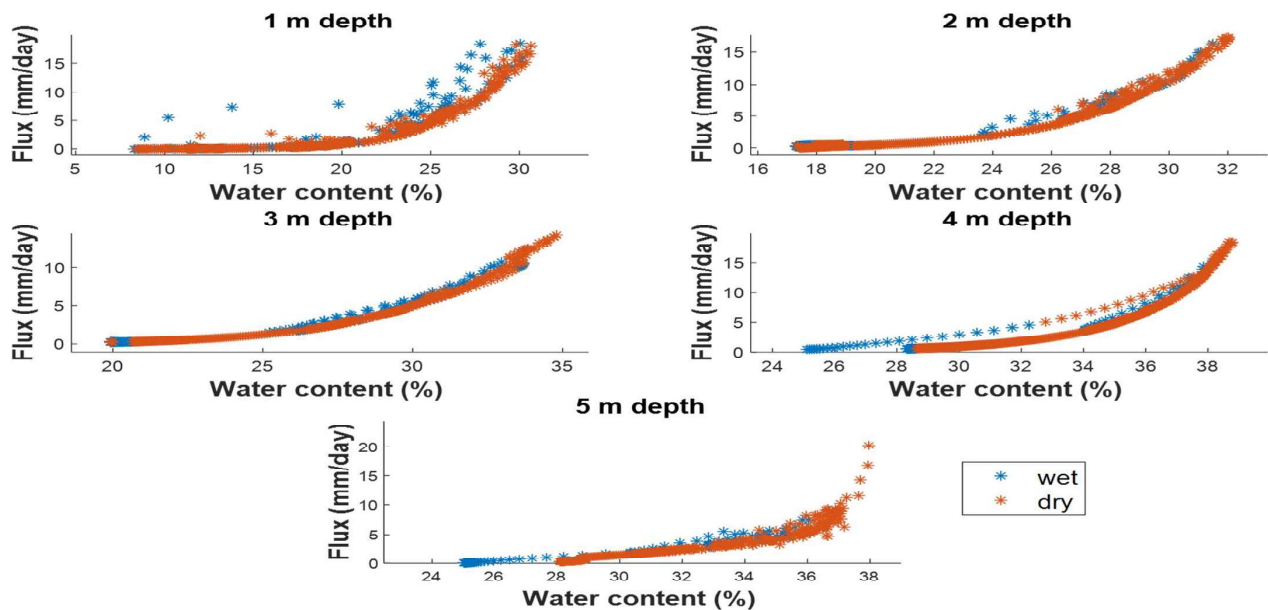


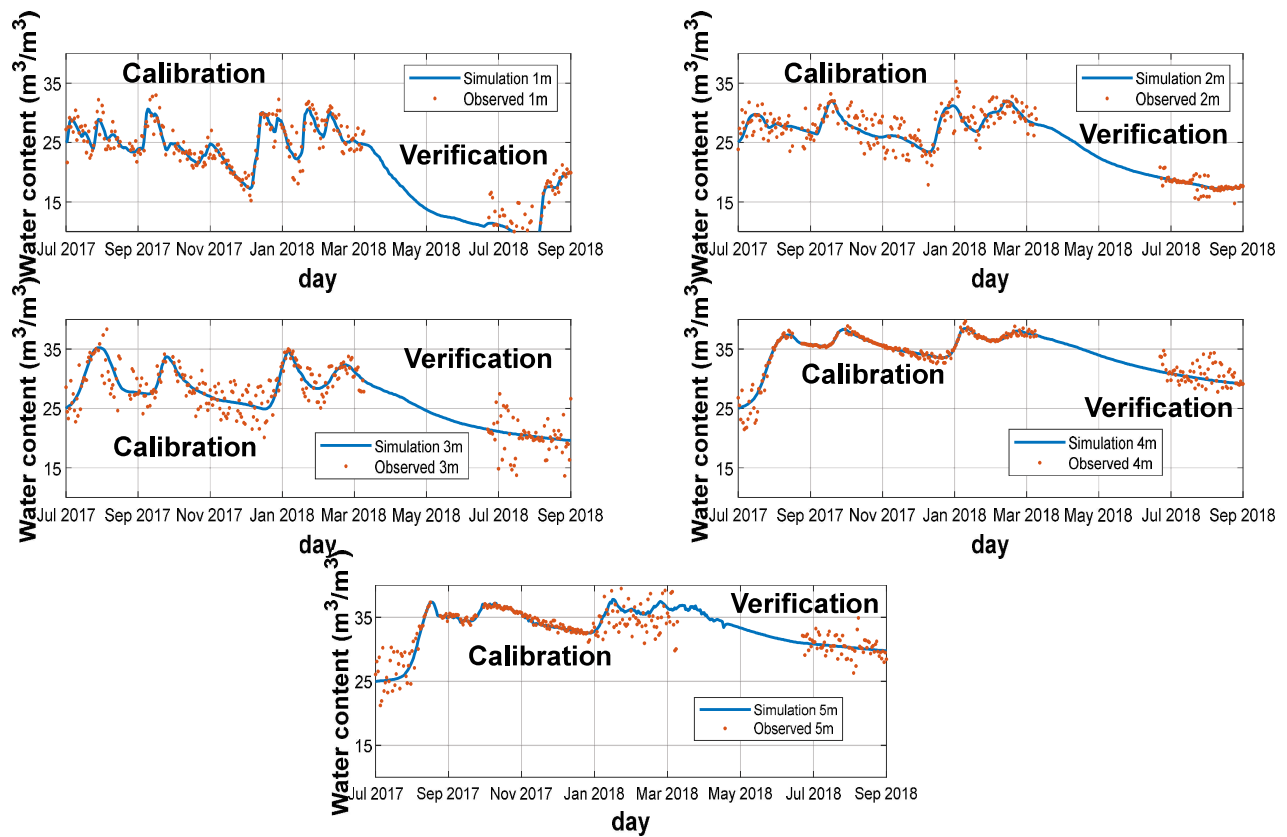
Source : Pwint phy aye, 2017

Results : Field measurement data



Relationship of flux and water content in the field test

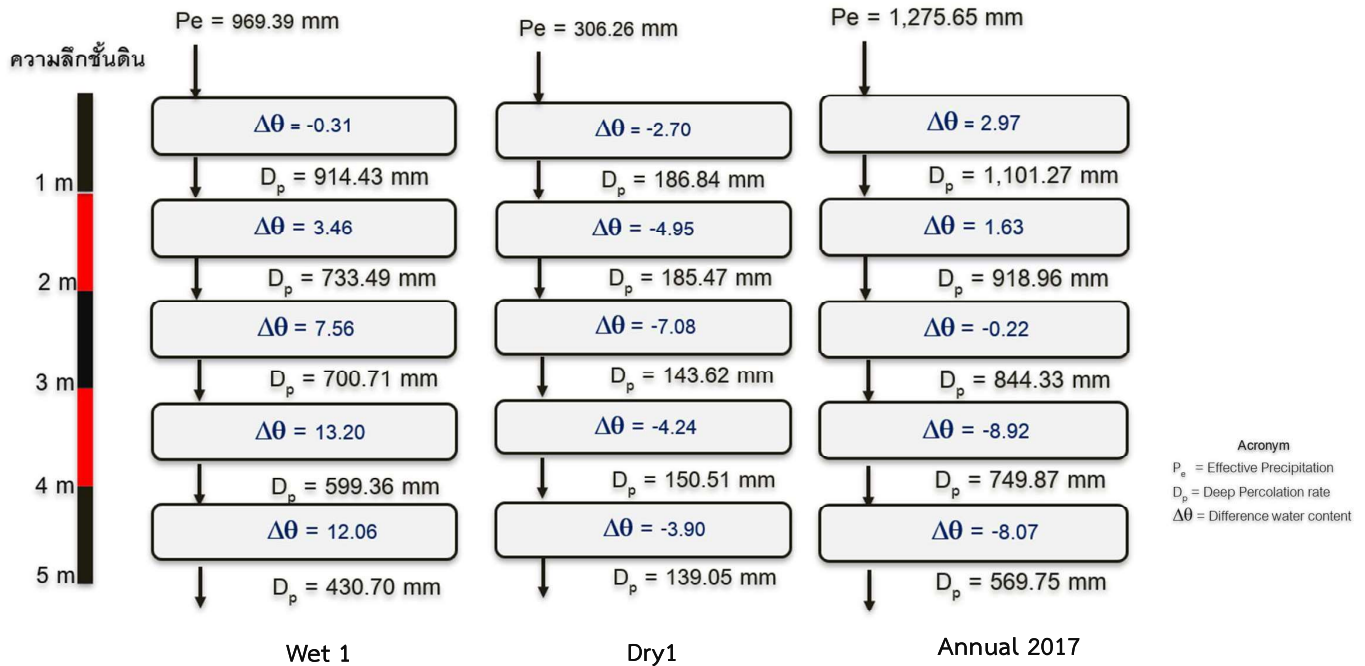




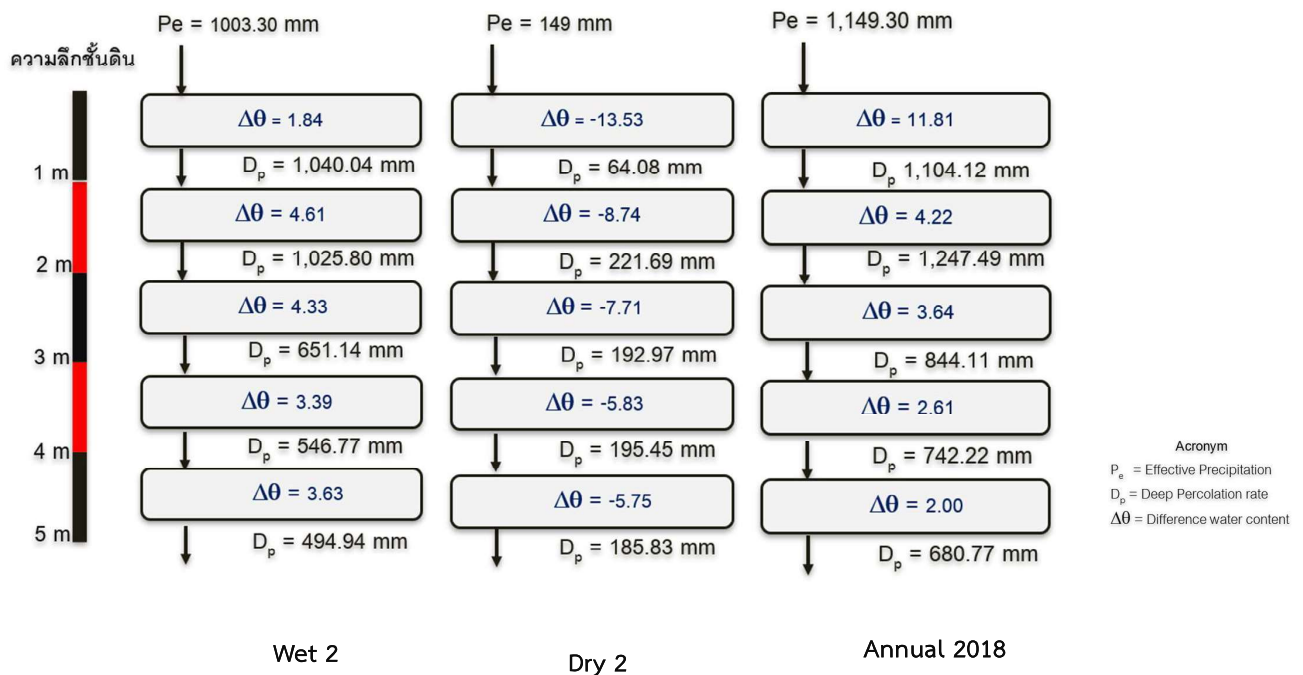
Soil retention parameters after calibration

Soil Depth (m)	Soil type	θ_r	θ_s	α	η	K_s
1	Sandy Loam	0.065	0.41	0.0075	1.89	170
2	Sandy Loam	0.065	0.41	0.0065	1.75	150
3	Sandy Loam	0.065	0.41	0.0045	1.65	90
4	Sandy Loam	0.065	0.41	0.0027	1.47	30
5	Sandy Clay	0.100	0.38	0.0021	1.35	25

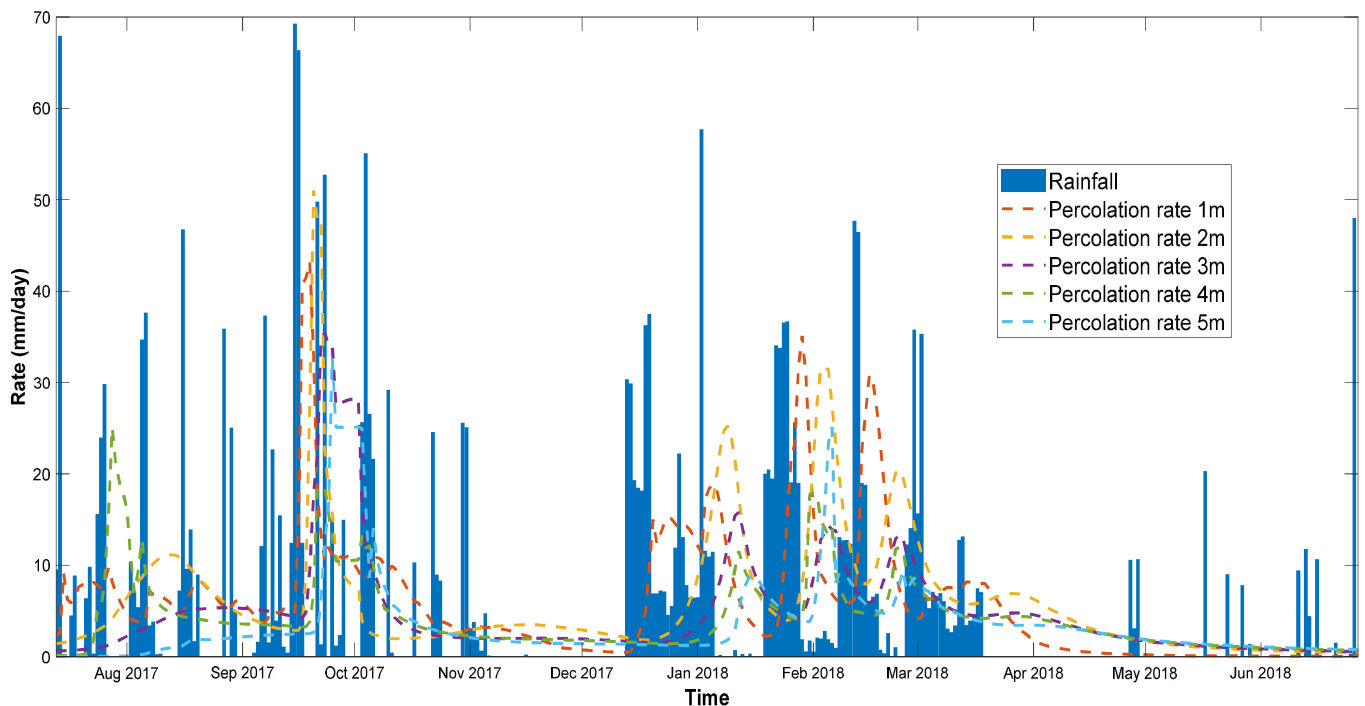
Water balance and water content change in 2017



Water balance and water content change in 2018



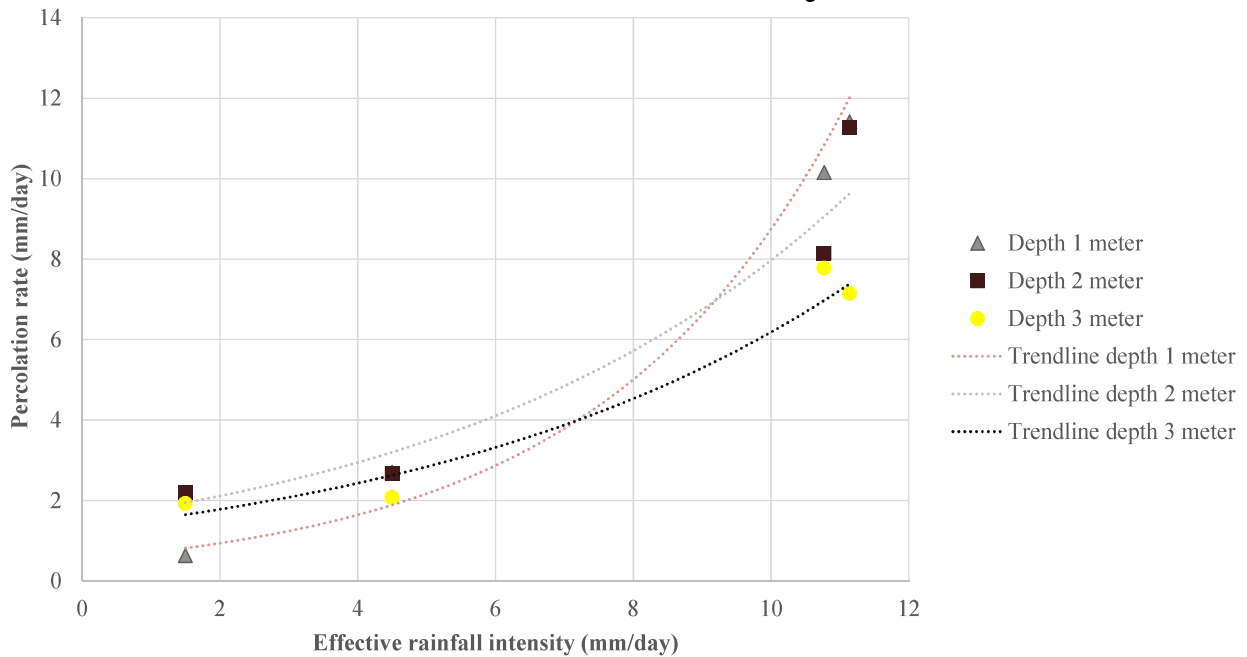
Percolation rate and Effective rainfall



Percolation rates at each soil depth and each period

Soil depth	Average percolation flux rate (mm /day)					
	Wet 1	Dry 1	Wet 2	Dry 2	Mean	
	July 13 – Oct 10, 2017 (90 days)	Oct 11- Dec 18, 2017 (69 days)	Dec 19, 2017 – Mar 19 , 2018 (91 days)	Mar 20 – June 27 2018 (100 days)	Wet (mm/day)	Dry (mm/day)
1 m	10.16	2.71	11.42	0.64	10.79	1.48
2 m	8.15	2.68	11.27	2.21	9.71	2.40
3 m	7.78	2.08	7.15	1.93	7.46	1.99
Effective rainfall	969.39	306.26	1003.3	149	10.96	3.00

The relationship percolation rate and Effective rainfall intensity



VI Conclusions

- ☐ The developed soil moisture sensor system gave reliable values of soil moisture under natural condition. The field monitored data with the application of HYDRUS 1-D model can estimate percolation rate and analyze water balance in each soil depth in more details.
- ☐ The percolation fluxes were compared with effective rainfall to find the percolation rate ratio in study area. The percolation rate ratios decrease from top to bottom depth. The highest ratio is sandy loam. The second ratio is sandy clay. The percolation flux of sand clay loam from top to bottom is 10.79 to 7.47 in wet period.

- ❑ The average rate of percolation to effective rainfall in each period is 0.43-0.49.
- ❑ The percolation rate is also affected from rainfall pattern and intensity, i.e., more intensified rainfall, more water content and more percolation, i.e., rainfall intensity of 2-11 mm/day induce percolation rate of 7-11 mm/day.
- ❑ The results from the model help to understand the deep percolation characteristics, the relationship of deep percolation rate with soil moisture content, effects of rainfall pattern and rainfall intensity.
- ❑ It also helped to estimate groundwater recharge and groundwater potentials in areas where the groundwater is used to mitigate the drought period.

Acknowledgements

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ANY QUESTIONS ?